Article

Open Access

SOURCE ROCK GEOCHEMISTRY OF THE PALEOGENE STRATA IN BENDE-UMUAHIA AND ITS ENVIRONS, NIGER DELTA BASIN, SOUTHEASTERN, NIGERIA

C. E. Oguanya¹, O. I. Chiaghanam², C. N. Nwokeabia³, K. C. Chiadikobi⁴ and N.O. Ikegwuonu⁵

^{1, 2, 4} Chukwuemeka Odumegwu Ojukwu University, Uli, Nigeria
³ Nnamdi Azikiwe University, Awka, Nigeria
⁵ Godfrey Okoye University, Ugwuomu-Nike, Nigeria

Received April 10, 2017; Accepted June 12, 2017

Abstract

Source rock geochemistry of the paleogene (Imo and Ameki Formations) strata in Bende- Umuahia and its environs, Niger Delta Basin, southern eastern Nigeria was assessed for its hydrocarbon play using walkley black wet oxidation and rock-eval pyrolysis techniques. The total organic carbon (TOC) average values for Imo and Ameki Formation are 0.86wt% and 0.74 wt%, this indicates a fair petroleum potential and organic matter concentration. The generative potential (GP) average value of 0.38mg Hc/g for the two formations studies indicates that the organic matter quality is poor. The result of hydrogen index (HI) shows a studied area that is basically type IV Kerogen. The average T_{max} value for Imo Formation was 408.8°C ad 410.8°C for Ameki Formation, which suggest that the studied area is thermally immature. The average vitrinite reflectance (RO) values for Imo and Ameki Formation, which are 0.59R₀ and 0.32R₀ respectively also suggests that the area studied is thermally immature. Theplots of Rock-Eval s2 versus Toc and hydrogen index versus oxygen index indicates that the studied area is basically type IV Kerogen, while the plots of production Index versus T_{max} and Hydrogen index versus T_{max} indicates thermally immature for the area. Source rock geochemistry of type IV Kerogen, Thermally immature and fair petroleum potential can be assigned to the area studied. *Keywords: Source Rock Geochemistry; Rock-Eval; Total Organic Carbon; Kerogen; Imo and Ameki Formation*.

1. Introduction

Niger Delta is a Cenozoic gross offlap clastic succession built out atop the Anambra Basin and forms part of the western African miogeocline that spread out onto the cooling and subsiding oceanic crust generated as the African and the South American lithospheric plates separated ^[1].

The systematic and consistent pattern of development of Niger Delta was such that, taken in a gross stratigraphic sense, three distinct vertically stacked lithologic units- Akata (Marine facies), Agbada (transitional, paralic facies) and Benin (Continental) have been, and are still been, constructed. The total sediment thickness is up to 12km beneath the upper and lower deltaic plain ^[1]. The outcropping units of the Cenozoic Niger Delta are Imo Formation (Paleocene-Early Eocene) which is blue –grey shales with sand lenses, marls, and fossiliferous limestones, sandstone members- Ebenebe, Umuna and Igbaku sandstones; shales with foraminifera and ostracods. Ameki group (Eocene- Early Oligocene) which comprises (a) Nsugbe Formation- mainly sands, with some conglomerate bands, (b) Nanka Formationcalcareous clays and silts with thin shelly limestones, rich in foraminifera (c) Ameki Formationcalcareous clays and silts with thin shelly limestone, rich in foraminifera, mainly sands, minor silt and clay intercalations. The Ogwashi- Asaba Formation (Oligocene- Miocene) comprises of clays, silts and sandstones with thin to thick lignite seams ^[1-3]. The geologic map of the study area is shown in Fig. 1. The objective of the present study is aimed at investigating the hydrocarbon potential of the outcropping shale units of the Cenozoic Niger Delta within the Umuahia-Bende axis which will help to:

- 1. Determine the quality of the source rock of the paleogene strata exposed in the Bende-Umuahia axis,
- 2. Determine and evaluate the thermal and hydrocarbon generation potential of the organic rich sediments of paleogene strata exposed in the Bende-Umuahia axis.

The deductions from the study will provide information necessary to optimize exploration activities in the Niger Delta Basin with a view of improving on past investigation/assessment in the study area.



Fig. 1. Geologic map of the study area

2. Geological setting and stratigraphy of the area

The cenozoic strata of the Niger Delta Basin are N-S to NW-SE striking units comprising Imo (Middle Paleocene- early Eocene), Ameki (Early to late Eocene), Ogwashi (Oligocene-Early Miocene), and Benin Formation (Miocene to recent); ^[4-5] (Fig. 2).



Fig. 2. The Cenozoic succession of the Niger Delta Basin (after [4-5])

The Imo Formation consist of thick blue to grey shales and mudstones, with occasional thin bands of calcareous sandstones and limestone, about 1000m thickness, ^[6]. The formation is exposed at Bende town, extending from Oduenyi village to Ndiwo junction, along Ndiwo-Ikporom road cut. This is overlain by Ameki Formation which consists of alternating shales, sandyshales, clay, sandstones, and fine-grained fossiliferous sandstone with thin bands of limestone ^[6-7]. At Itumbuzor in Bende, the formation directly overlies the Imo Formation and terminates westward around Isiadu Ameke and Amaogugu where it is completely buried by the overlying Ogwashi Formation ^[4,8]. The Ogwashi Formation ^[9], consists of alternating coarse grained sandstones, light coloured clays of continental Origin, with Lignite seams. The formation is overlain by the alluvial deposits of Benin Formation. These Cenozoic succession up-dip of the Niger- Delta have been correlated to the subsurface down-dip formations ^[3,10] Fig. 3, and since there are no stratigraphic breaks, the formation in the up-dip and down-dip constitute the Niger delta basin ^[4], according to the rules of stratigraphic nomenclature ^[2,6,11] Table 1.



Fig 3. Stratigraphic equivalencies between the outcropping and the subsurface Niger Delta (*after* ^[3])

Table 1. Stratigraphic column of Niger Delta Basin (*modified after Frankyl and Cordry* ^[10])

3. Previous work

Source rock geochemistry studies in the Niger Delta Basin and Late Cretaceous Anambra Basin has the undertaken by various workers, these workers includes Chiaghanam et al. [12], studies the palynofacies and kerogen Analysis of Upper Cretaceous (Early-Campanian to Maastrictian) Enugu Shale and Mamu Formation in Anambra Basin, Southeastern, Nigeria. The source rock potential and thermal maturity of the Eocene Nanka Formation in Anambra Basin: An Appraisal of Ogbunike reference locality, SE, Nigeria was carried out by Chiaghanam et al. [13]. Palynology, source rock potential and thermal maturity of Eocene Nanka Formation in Anambra basin: An investigation of Agulu Lake, S.E. Nigeria ^[14]. Source rock and thermal maturation of Campanian Enuqu Shale in Anambra Basin, S.E. Nigeria. Anozie et al. ^[15]. The Depositional environments, Organic richness and petroleum generating potential of the Campanian to Maastrictitian Enuqu Formation were done by Ojo et al. [16]. Ogala [17] studies hydrocarbon potential of the Upper Cretaceous coal and shale units in the Anambra Basin. source rock maturation studies using vitrinite reflectance and geothermal data from six wells in Gaba and Wabi, fields, onshore Niger delta, was done by Didei and Akana ^[18]. Geochemical investigation of potential source rocks for Agbada Formation, Osioka South Area, Western Niger Delta was studies by Nyantakyi et al. ^[19]. Source Rock characterization of Agbada Formation in well Z, offshore, Niger Delta Asadu et al. [20].

The present study attempts to provide information on the hydrocarbon potential and thermal maturation of the study area, since little or no work has been done on its source rock geochemistry.

4. Materials and methods

Ten shale samples recovered from the outcropped Imo and Ameki Formation in the study are selected for Total organic content determination and Rock eval pyrolysis which was aimed towards: a) Determining the quantity, or amount of organic matter.

- b) Determining the quality, or type of organic matter
- c) Determining the thermal maturity, or extent of burial/heating ^[21].

Total Organic Carbon (TOC): The ten shale samples were subjected to preliminary total organic carbon content determination by using Walkley Back wet oxidation method, which involves subjecting 0.5g of each pulverized samples to chromic oxidation using the principles of Walkley and Black ^[22]. This served as a preliminary step towards carrying out rock- eval pyrolysis analysis.

Rock-eval Pyrolysis: The samples were subjected to rock-eval pyrolysis using the principles and procedures according to Espitalie *et al.* ^[22] and Stach *et. al.* ^[23]. The samples were heated in an inert atmosphere to 550°C using a special temperature programme. The parameters/results obtained from the analysis include;

- 1) First Peak ((S_1) hydrocarbon generated at the temperature of 300°C.
- 2) Second Peak (S₂) hydrocarbon yield from cracking of Kerogen and heavy kerogen in the rock sample at the temperature of range of 300-550°C.
- 3) Third Peak (S₃) carbon dioxide (CO₂) generated during the process of thermal cracking of kerogen.
- 4) T_{max} which is the value of thermal maturity and corresponds to the Rock eval pyrolysis oven temperature (°C) at maximum S₂ generation ^[25].
- 5) Production or productivity index $[PI = S_1/(S_1 + S_2)]$.
- 6) Hydrogen index [HI = (S₂/Toc) x 100.mg HC /gToc]
- 7) Oxygen index [OI = $(S_3/Toc] \times 100.mg CO_2/g Toc]$.
- 8) Calculated vitrinite reflectance (R_o).

5. Results and discussion

Research has shown the importance of organic geochemical methods in assessing the generative potential and characteristics of source rocks ^[15,21,24-27]. In this work, the petroleum potential (quantity), Kerogen type (quality) and level of thermal maturity of the analysed samples of the paleogene (Imo and Ameki Formation) were discussed based on rock-eval pyroysis and total organic carbon.

Organic Matter Richness: Total organic carbon content (TOC) and Rock-eval analysis were carried on 10 shale samples that are presumed to be source rocks with 5 samples each from Imo Formation and Ameki Formation. The Total organic carbon in a source rock comprises three base components. (a) organic carbon in a retained hydrocarbons as received in the laboratory, (b) organic carbon that can be converted to hydrocarbon, called convertible carbon ^[28] or reactive or labile carbon ^[29]; and (c) carbonaceous organic residues that will not yield hydrocarbon because of insufficient hydrogen commonly referred to as inert carbon ^[15,17,28-29].

The total organic carbon (Toc) ranges from 0.18 to 1.43 wt% with an average value of 0.86 wt% for Imo Formation, and 0.40 to 1.22 wt% with an average value of 0.74 wt% for Ameki Formation (Table 2a and 2b). This indicates a fair Petroleum potential and Organic matter concentration for the studied area ^[21,30-31].

The quality of the source rock in the studied area is confirmed by the pyrolysis -derived generative potential (G.P. = $S_1 + S_2$). The rock-eval pyrolysis revealed that the total hydrocarbon generative potential ranges from 0.17 to 0.63 mg HC/g rock with an average value of 0.38 Mg HC/g rock for Imo Formation and 0.19 to 0.55 mg HC/g rock with as average valve of 0.38 mg HC/g rock for Ameki Formation. This indicates that the organic matter quality for the studied area is poor. S_1 measures hydrocarbon shows as the amount of free hydrocarbon

that can be volatilized out of the rock without cracking the kerogen (mg HC/g rock). S_1 increases at the expense of S_2 with maturity while S_2 measures the hydrocarbon yield from cracking of kerogen (mg Hc/g rock) and heavy hydrocarbon and represents the existing potential of a rock to generate petroleum ^[21].

Station Name	Station no.	Sample Type	Formation	TOC	Rock-Eval			Tmax (°C)
					S1	S2	S3	
OKWANKA STREAM	L1	OUTCROP	IMO	1.34	0.05	0.54	0.56	432
INYIAKWU STREAM	L2	OUTCROP	IMO	0.54	0.07	0.14	0.29	416
OKOTU JUNCTION	L4	OUTCROP	IMO	0.18	0.06	0.11	0.25	344
NDIWO/IKPOROM ROAD	L6	OUTCROP	ΙΜΟ	1.43	0.11	0.17	0.44	438
NDIWO JUNCTION	L8	OUTCROP	IMO	0.82	0.13	0.37	0.42	414
				0.86	0.08	0.27	0.39	408.8

Table 2a. Results of geochemical analysis for Imo Formation

	Ro	HI	OI	S2/S3	S1/TOC	PI	GP S1+S2
OKWANKA STREAM	0.62	40	42	1.0	4	0.08	0.59
INYIAKWU STREAM	0.33	26	54	0.5	13	0.33	0.21
OKOTU JUNCTION	0.97	60	137	0.4	33	0.35	0.17
NDIWO/IKPOROM ROAD	0.72	40	103	0.4	26	0.39	0.28
NDIWO JUNCTION	0.29	45	52	0.9	16	0.26	0.63
	0.59	42.2	77.6	0.64	18.4	0.28	0.38

Table 2b. Result of geochemical analysis for Ameki Formation

Station Name	Station no.	Sample Type	Formation	тос	Rock-Eval			Tmax (°C)	
					S1	S2	S3		
OKWANKA STREAM	L1	OUTCROP	Ameki	0.54	0.08	0.47	1.22	429	
INYIAKWU STREAM OKOTU JUNCTION	L2 L4	OUTCROP OUTCROP	Ameki Ameki	0.75 0.71	0.04 0.07	0.37 0.49	0.35 1.11	415 413	
NDIWO/IKPOROM ROAD	L6	OUTCROP	Ameki	1.22	0.05	0.22	1.02	386	
NDIWO JUNCTION	L8	OUTCROP	Ameki	0.40	0.08	0.11	1.31	411	
				0.74	0.06	0.33	1.00	410.8	
	Ro	HI	OI	S2/S3	S1/1	ГОС	PI	GP S1+S	2
OKWANKA STREAM	0.56	87	227	0.4		15	0.15	0.5	55
INYIAKWU STREAM	0.31	49	47	1.1		5	0.10	0.4	¥1
OKOTU JUNCTION	0.27	69	157	0.4	10		0.13	0.5	50
NDIWO/IKPOROM ROAD	0.21	18	84	0.2	4		0.19	0.2	27
NDIWO JUNCTION	0.24	28	329	0.1		20	0.42	0.1	19
	0.32	50.2	168.8	0.44	1	0.8	0.20	0.3	38

The hydrogen index (HI) values range from 26 to 60mg Hc/g Toc, with an average value of 42.2 mg HC/g Toc for Imo Formation, and 18 to 87 mg HC/g Toc, with average of 50.2HC/g TOC for Ameki Formation. This indicates a source rock that cannot expelled oil or gas at peak

maturity and a type IV Kerogen, with atomic H/C of < 0.7 and S_2/S_3 of <1 ^[21,32], however, the average value for Ameki Formation falls at the boundary and may be regarded as type III Kerogen.

Type of organic matter



The types of organic matter in sediments of the studied area were analyzed using rock-eval pyrolysis. The ten shale samples assessed by rock-eval pyrolysis are basically of type IV Kerogen, which those not expel oil or gas at the peak of maturity and can be regarded as being inert. The plot of Rock-eval S₂ versus Toc (Figure 4) describes the hydrocarbon potential of the samples. The plot reveals that the samples analyzed have minimum value required for quality source rock. The relationship between the hydrogen index (HI) versus oxygen index (OI) (Figure 5) suggests a Kerogen Type IV organic matter which are predominantly inert.

Fig 4. Plot of Rock-eval S₂ versus TOC

Thermal maturity of organic matter

Thermal maturity or extent of burial heating provides an indication of the maximum paleo temperature reached by a source rock ^[17]. The thermal maturation of Niger Delta Basin has been studied by authors ^[18-20]. The degree of Thermal maturity of Paleogene strata of the



Fig 5. Plot of Hydrogen Index versus Oxygen Index

studies area was assessed using pyrolysisparameters derived such as vitrinite reflectance (Ro), production index, Rockeval and T_{max} which measures thermal maturity and corresponds to the rock eval, pyrolysis oven temperature (°C) at maximum S₂ generation ^[21]. The T_{max} value represents the temperature at which the largest amount of hydrocarbon is produced in the laboratory when a whole rock sample undergoes a pyrolysis treatment ^[15]. The T_{max} value in the studied area ranges.

From 344° C to 438° C with an average valve of 408.8° C for Imo Formation, while 386° C to 429° C with an average valve 410.8° C was recorded for Amaki Formation. This indicates that the studies samples stage of thermal maturity for oil is still immature. The calculated vitrinite reflectance (R_o) value which increases during thermal maturation due to complex, irreversible aromatization reaction ^[21], ranges from 0.29 to 0.97 R_o with an average valve of 0.59 R_o for Imo Formation and ranges from 0.21 to 0.56 R_o



Fig. 6: Plot of production index (PI) against maturity (Based on Tmax)

with an average valve of 0.32 R_0 for Ameki Formation. The average valve for R_0 in the studied area suggests that stage of thermal maturity for oil is immature. The result of vitrinite reflectance (R_0) in the studies area shows that the Thermal alteration index will range between 1.5 to 2.6. Thermal alteration index (TAI) is a numerical scale based on thermally induced colour changes in spores and pollen [²¹].

The Plot of production index (PI) against maturity (Based on Tmax) as shown in Fig 6. and plot of hydrogen index [HI (S2/Toc x 100, mg/g Toc] against maturity (Tmax) as shown in Fig 7. Suggests that the studied area are basically type IV Kerogen, meaning a "dead

carbon" showing low atomic H/C (about 0.5-0.6) and low to high O/C (\leq 0.3). It also shows that the studied area are basically immature (level of thermal maturation) except for sample location L 6 (Ndiwo/ Ikpeoram road) that falls within the thermally matured /oil window range.



Fig 7. Plot of Hydrogen Index versus T_{max}

6. Conclusion

This study has shown that the shale of peleogene strata (Imo & Ameki Formations) in the Bende Umuahia and its environs in the Niger-Delta Basin, South-Eastern Nigeria has a Total

Organic Carbon whose petroleum potential is fair, with poor organic matter quality. The hydrogen index value that is basically type IV Kerogen. The T_{max} and Vitrinite reflectance results all indicates that the studied area is thermally immature, except for L6 which can be regard as being a type III. kerogen and thermally matured.

References

- [1] Nwajide CS. 2013. Geology of Nigeria's Sedimentary Basins. CSS Bookshops Limited 50/52 Broad Street, Lagos, Nigeria 548P.
- [2] Maron P. Stratigraphical Aspects of the Niger Delta. J. Min and Geol., 1969; 4(1–2): 3-12.
- [3] Wright JB, Hastings DA, Jones WB and Williams HR. 1985. Geology and Mineral Resources of West Africa. George Allen and Unwin, London, 187P.
- [4] Ikegwuonu ON and Umeji OP. Palynological age and paleoenvironment of deposition of Mid-Cenozoic sediments around Umuahia Niger Delta, Southeastern Nigeria. Journal of African Earth Sciences 2016; 117: 160-170.
- [5] Murat RC 1972, Stratigraphy and Paleography of the Cretaceous and Lower Tertiary in Southern Nigeria. In Whiteman, Dessalvagia (eds), Conference of African Geology. Ibadan University Press. Pp. 251-266.
- [6] Reyment RA. 1965. Aspects of the Geology of Nigeria. University Press, Ibadan, 144 pp.
- [7] Arua I. Paleoenvironment of Eocene Ameki Formation of the Afikpo Syncline, Southern Nigeria J. African Earth Sci.1986; 5: 279 294.
- [8] Ikegwuonu, ON. 2015. Middle Paleocene to Early Miocene Palynostratigraphy of Sediments in Bende - Umuahia Area, Niger Delta, Basin, Southern Nigeria. Unpublished M.Sc. Thesis, Department of Geology, University of Nigeria Nsukka, pp 205.
- [9] Kogbe CA. 1976. The Cretaceous and Paleogene sediments of Southern Nigeria in Kogbe, C.A. (Ed.), Geology of Nigeria. Elizabethan publishing company, Ibadan, Nigeria pp. 237-252.
- [10] Frankyl EJ and Cordry E A. 1967. The Niger Delta Oil Province- recent developments onshore and offshore. Proc. 7th World Pet. Congr. Mexico City IB, pp. 195-209.
- [11 Short KC and Stauble AJ. Online of Geology of Niger Delta. Bulletin of American Association of Petroleum Geologists, 1967; 51(5): 761-779.
- [12] Chiaghanam OI, Chiadikobi KC, Ikegwuonwu ON, Omoboriowo AO, Onyemesil, OC and Acra EJ. Palynofacies and Kerogen Analysis of Upper cretaceous (Early Campanian to Maastrictian) Enugu Shale and Mamu Formation in Anambra Basin, South Eastern, Nigeria. International Journal of Scientific and Technology research, 2013; 2(7): 225-229.
- [13] Chiaghanam OI, Chiadikobi KC, Ikegwuonwu ON and Omoboriowo AO. Palynology, source rock potential and thermal maturity of Eocene Namka formation (Ameki group) in Anambra Basin: An investigation of Agulu Lake. South eastern Nigeria 10SR-Journal of Applied Geology and geophysics, 2014; 2(5): 87-97.
- [15] Anozie HC, Chiaghanam OI and Chiadikobi KC. Source Rock and Thermal Maturation of Campanian Enugu Shale in Anambra Basin, Southeastern, Nigeria. International Journal of Geology, Earth and Environment sciences, 2015; 4(3): 296-307.
- [16] Ojo OJ, Kolawole AU and Akande SO. Depositional Environments, Organic richness and Petroleum Generating Potential of the companion to maastriafitian Enugu Formation. Anambra Basin, Nigeria. The Pacific Journal of Science and Technology, 2009; 10(1): 614 -628.
- [17] Ogala JE. Hydrocarbon Potential of the Upper Cretaceous coal and shale units in the Anambra Basin, South Eastern Nigeria. Pet Coal. 2011; 53 (1): 35-44.
- [18] Didei IS and Akana TS. Source Rock Maturation studies using vitrinite reflectance and geothermal data from six wells in Gabo and Wabi fields, onshore Niger Delta, Nigeria International Journal Geology and mining. 2016; 2(2): 064-070.
- [19] Nyantakyi EK, Hu WS, Borkloe JK Nagre RD and Frimpong KI. Geochemical investigation of potential source rocks for Agbada formation, Osioka South Area, Western Niger Delta, Nigeria geosciences, 2014; 4(1): 13-22.
- [20] Asadu AN, Omo-Irabor OO and Ibe KA. Source Rock Characterization of Agbada Formation in Well Z, Offshore Niger Delta, Nigeria. International Journal for Research in Emerging Science and technology. 2015; 2(6): 92-100.
- [21] Peters KE and Cassa MR. 1994. Applied Source Rock Geochemistry in Magoon, L.B and W.G. Dow, eds. 1994. The Petroleum system –from source to trap. AAPG memoir 60.

- [22] Espitalie J. Marquis F and Barsony, I. 1984. Geochemical Logging. In Analytical Pyrolysis-Techniques and Applications. Vootrees, K.J. (editor). Butterworth. Guildford, U.K. Pp 276-304.
- [22] Walkley A and Black TA. An Examination of the Titration method for determining soil organic matter and proposed modification of the chromic acid titration methods. Soil science. 1965; 37: 29-38.
- [23] Stach E, Mackowsky MT, Taylor GH, Chandra D, Teichmoller MI and Teichmuller R. 1982. Stach's Textbook of Coal Petrology, 3rd Edition, Stuttgart, Borntraeger, 535P.
- [24] Baskin DK. Atomic H/C Ratio of Kerogen as an Estimate of Thermal maturity and organic matter conversion AAPG. Bull. 1997; 81: 1437-1450.
- [25] Peters .E. Guideline for evaluating Petroleum source Rocks using programmed Analysis AAPE Bulletin. 1986; 70: 318-329.
- [26] Akande SO, Ojo OJ, Erdtmann BD and Hetenyi M. 1998. Paleoenvironments, Source Rock Potential and Thermal maturing of the upper Benue Rift Basins. Nigeria: Implications for hydrocaborn explortation" Organic Geochemistry 1998; 29: 531 -542.
- [27] Akande SO. Paleoenvironment, Organic Petrology and Rock-oval studies on source rock facies of the Lower Maastrichtian, Patti Formation, Southern Bida Basin, Nigeria. Journal of African earth Sciences, 2002; 25:1233-1266.
- [28] Jarvie DM. 1991. Total Organic carbon (TOC) Analysis in Treatise of Petroleum Geology. Handbook of petroleum geology, source and migration processes and evaluation techniques edited by Merril R.K. AAPG bull. Pp 113-118.
- [29] Cooles GP, Mackezie AS and Quigley TM. 1986. Calculation of petroleum masses generated and expelled from source rocks. In Advances in Organic geochemistry, Oxford, Pergamon, Organic Geochemistry, edited by Leythaeuser. D. and Rullkotter, J. 10 pp. 235-245.
- [30] Hedberg HD and Moody JO. Petroleum prospects of Deep offshore. AAPG Bull., 1979; 63: 286-300.
- [31] Hunt JM .1979. Petroleum Geochemistry and Geology, 2nd ed. Freeman and Company. San Francisco. Pp743.
- [32] Tissot B and Welte DH. 1984. Petroleum Formation and occurrence. Springer-Verlag: New York, NY. PP. 699.
- [33] Chiaghanam OI, Chiadikobi KC, Ikegwuonwu ON, Omoboriowo AO, Onyemesili OC and Yikarebogha Y. Source Rock potential and Thermal maturity of the Eocene of Nanka Formation (Ameki Group) in Anambra Basin: An appraisal of Ogbunike Reference locality. South Eastern, Nigeria. 10SR-Journal of Applied Geology and Geophysics. 2014; 2(3): 11-17.

To whom correspondence should be addressed: oichiaghanam@yahoo.com